

# *Use of data assimilation methods in coastal sediment models*

Conference or Workshop Item

Presentation

Dance, S. L., Baines, M. J., Lawless, A. S., Mason, D. C., Nichols, N. K., Scott, T. R., Sweby, P. K., Smith, P. and Thornhill, G. D. (2010) Use of data assimilation methods in coastal sediment models. In: Flood Risk from Extreme Events (FREE) Final Science Meeting, 20 October, London. (Unpublished) Available at <http://centaur.reading.ac.uk/8213/>

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# Use of data assimilation methods in coastal sediment models

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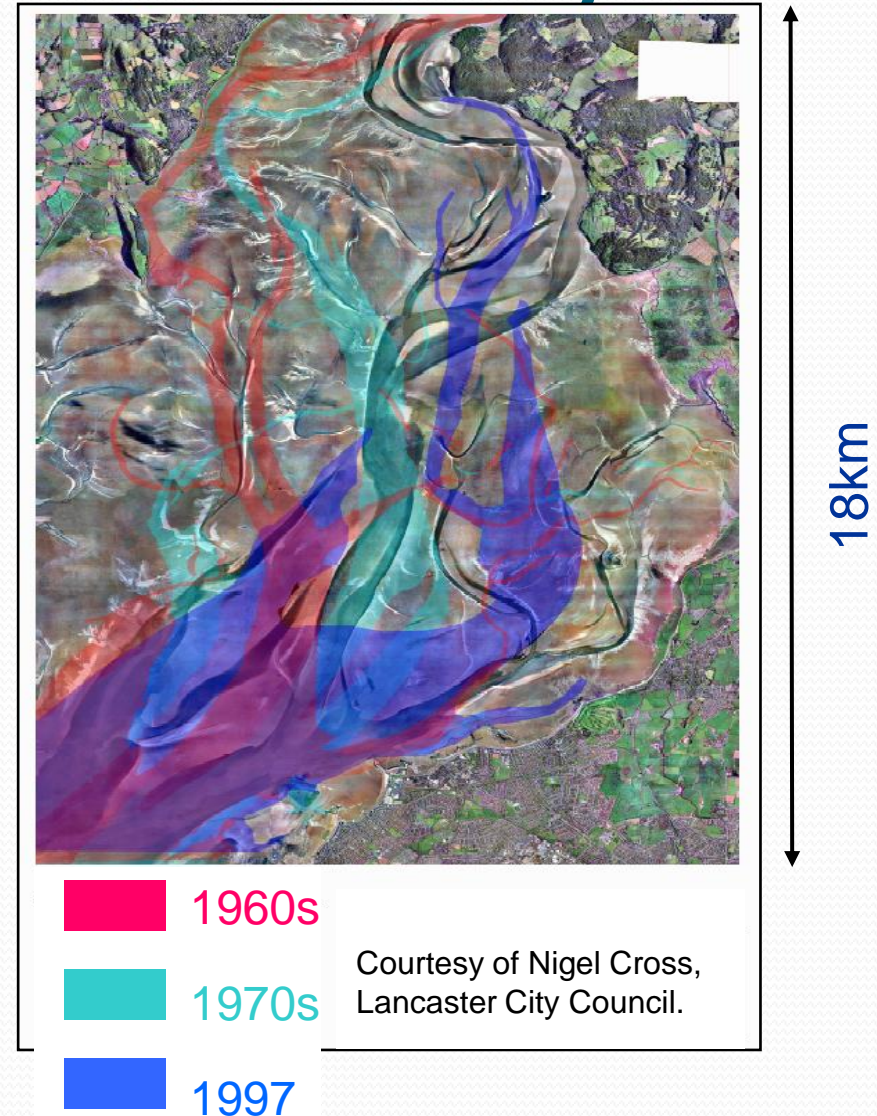
With thanks to Kevin Horsburgh, POL and the Environment Agency

# Outline

- Motivation: morphodynamics and flooding
- Observational data and morphodynamic model
- Data assimilation
- Parameter estimation
- Conclusions

# Case study: Morecambe Bay

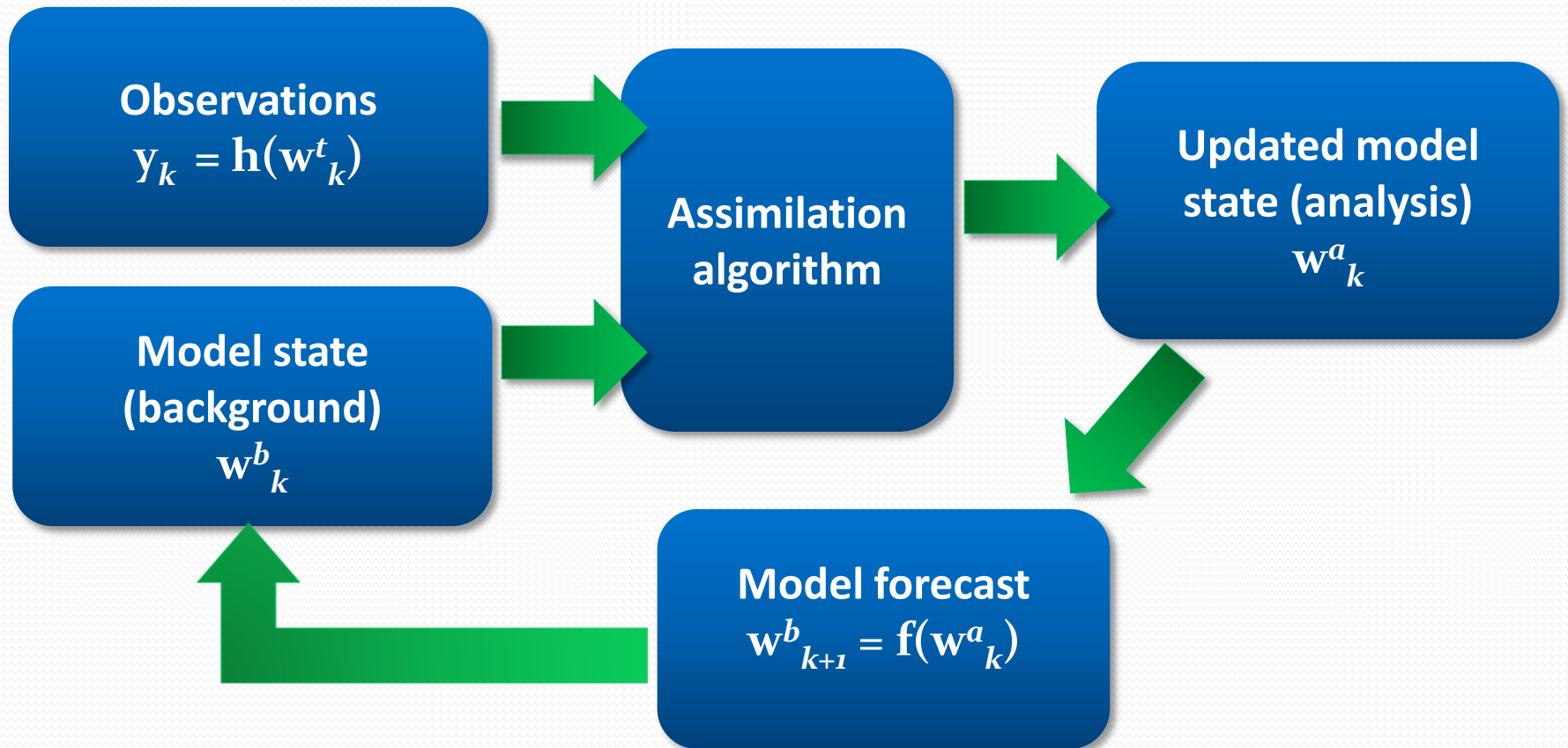
- Channels in Morecambe Bay can move by several kilometres in just a few years
- Channel movement
  - impacts on habitats in the bay
  - affects access to ports
  - has implications for flooding during storm events
- For example, Morecambe can be flooded by storm waves propagating up the deep-water channels



# The general idea

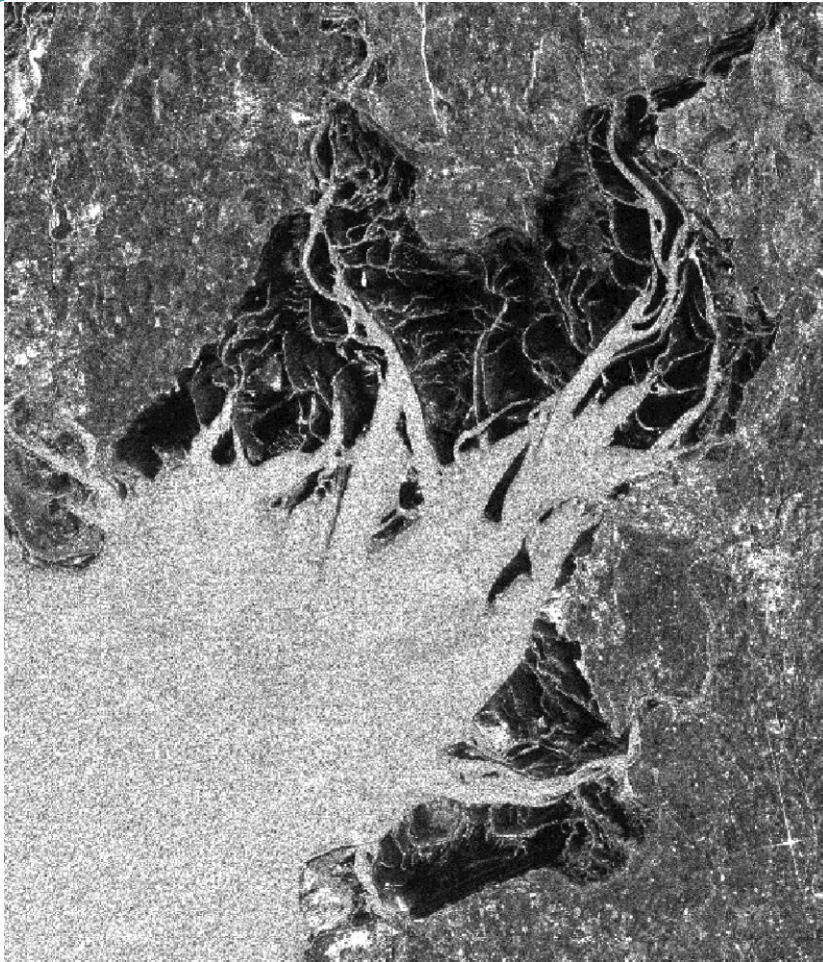
- Coastal bathymetry is dynamic and evolves with time
  - water action erodes, transports, and deposits sediment, which changes the bathymetry, which alters the water action, and so on
- Accurate knowledge of coastal bathymetry at the time of a storm event would allow improved flood forecasting using coastal flood inundation models
  - but it is impractical to continually monitor large coastal areas in anticipation of a storm
- A solution may be to run an operational coastal area morphodynamic model
  - and keep the model on track using data assimilation
- As observations become available they can be used to nudge model bathymetry back towards true bathymetry
  - these observations may be infrequent and only sample a small part of the model domain
- Data assimilation can also be used for parameter estimation

# Sequential data assimilation

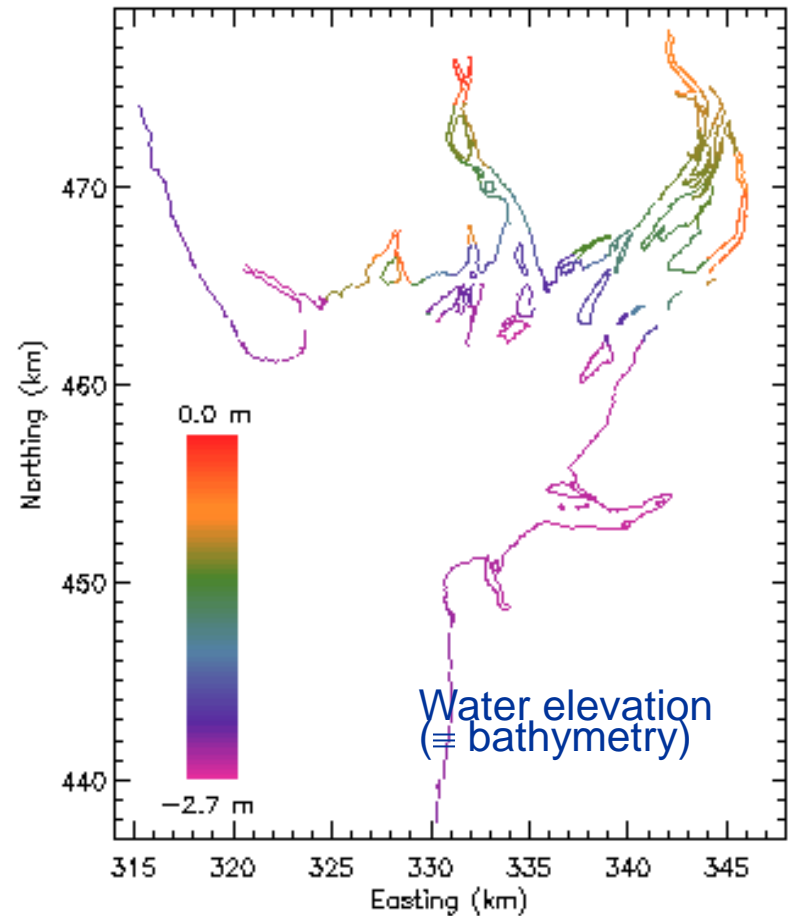




# Waterline data for Morecambe Bay

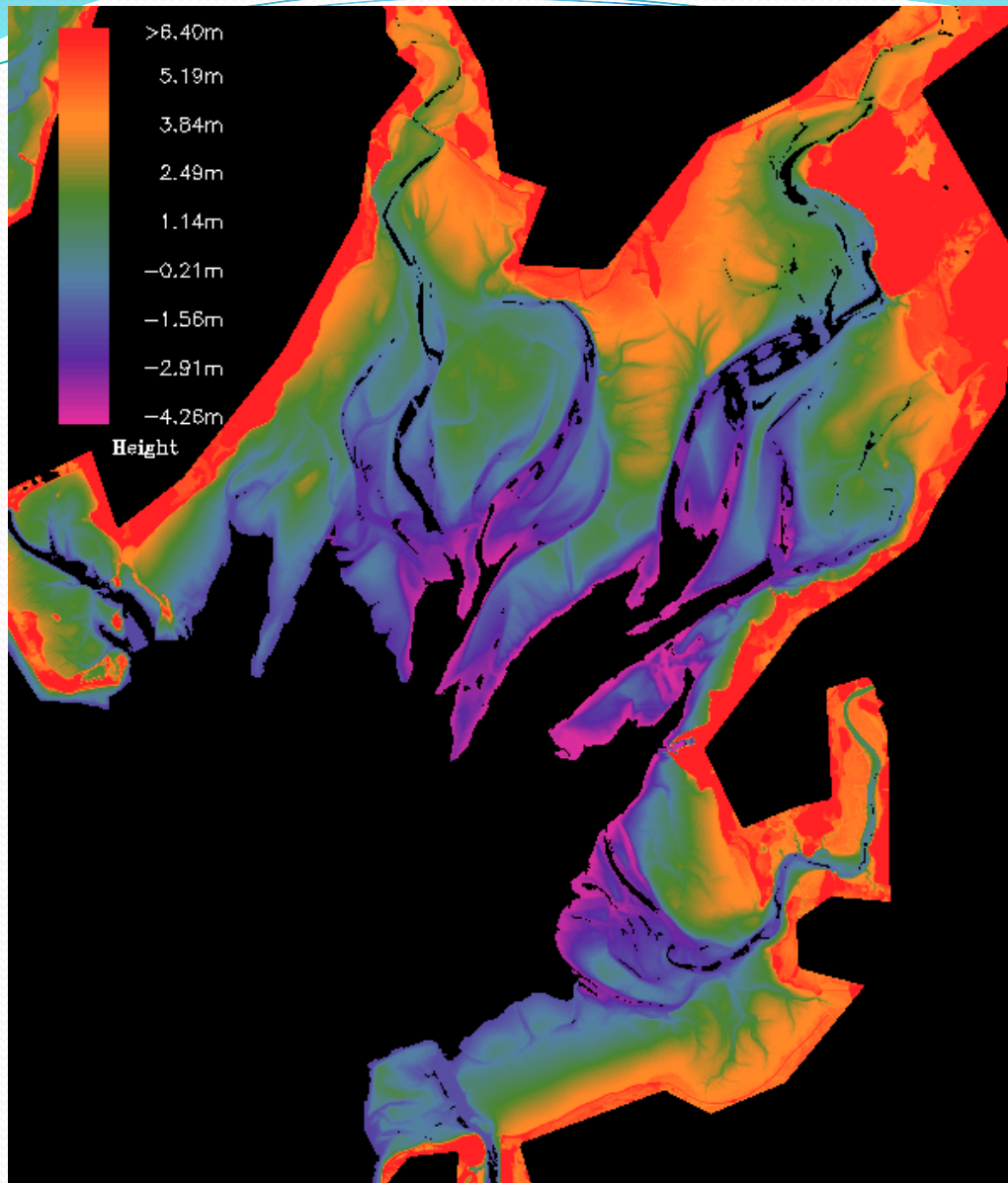


An image of the land-sea boundary  
(satellite SAR image © ESA)



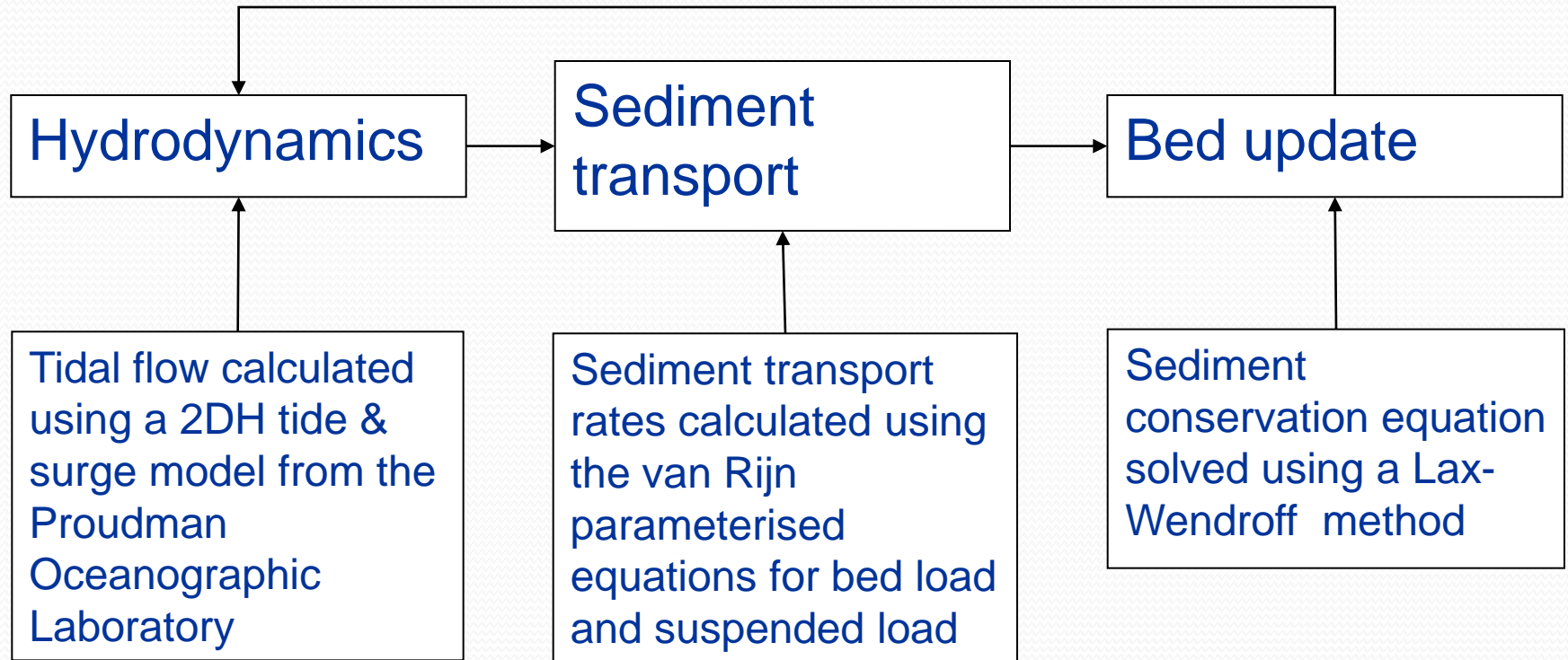
Water elevation predicted along the boundary by a hydrodynamic model with surge component





*LiDAR data from  
15/11/05 to be  
used for  
validation.*

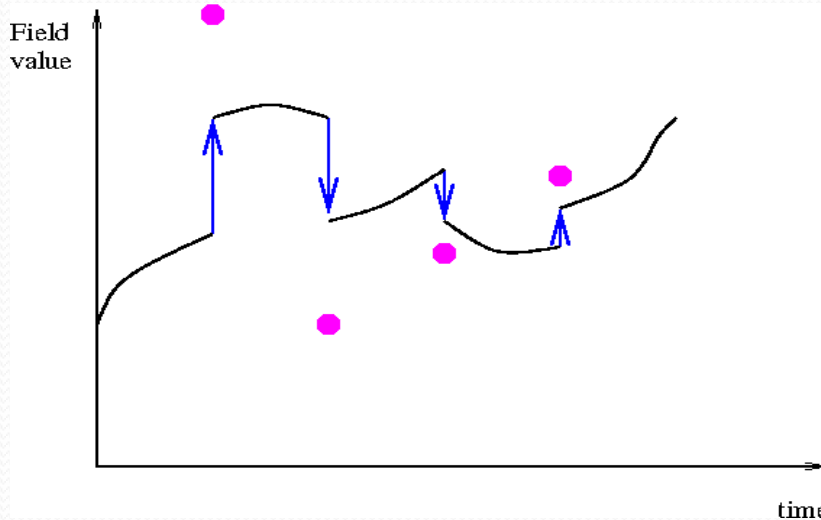
# Morphodynamic model



Spatial resolution ~ 240m

A simple model, relative to state-of-the-art engineering models, but adequate for assessing the benefits of data assimilation

# 3D Var data assimilation

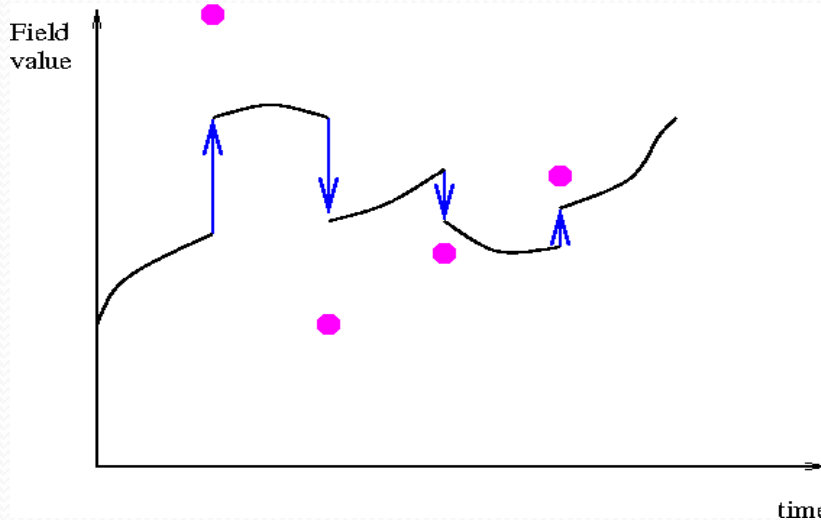


- Analysis state is found by minimising a cost function

$$J(\mathbf{w}_k) = \underbrace{(\mathbf{w}_k - \mathbf{w}_k^b)^T \mathbf{B}_k^{-1} (\mathbf{w}_k - \mathbf{w}_k^b)}_{\text{background term}} + \underbrace{(\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))^T \mathbf{R}_k^{-1} (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))}_{\text{observation term}}$$

$\mathbf{B}$  and  $\mathbf{R}$  are the covariance matrices of the background and observation errors

# 3D Var data assimilation



- Analysis state is found by minimising a cost function

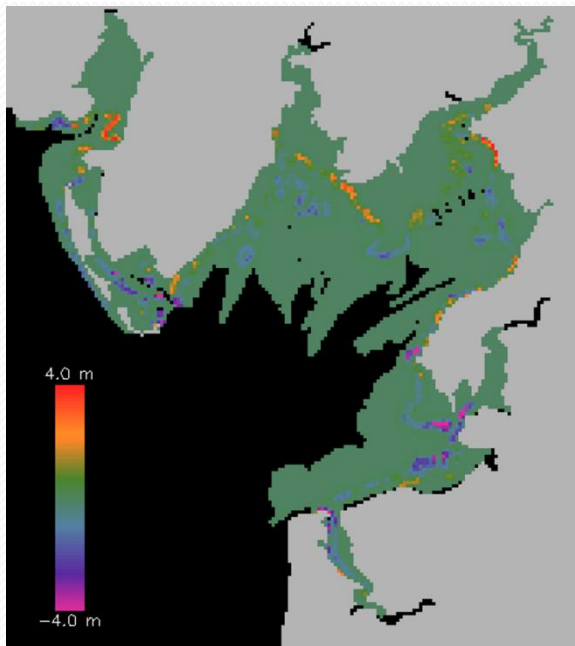
$$J(\mathbf{w}_k) = (\mathbf{w}_k - \mathbf{w}_k^b)^T \mathbf{B}_k^{-1} (\mathbf{w}_k - \mathbf{w}_k^b) + (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))^T \mathbf{R}_k^{-1} (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))$$

background term      observation term

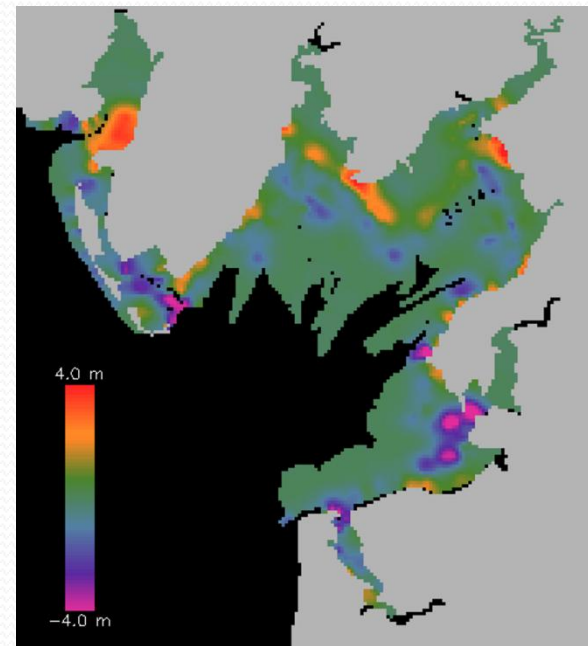
**B** and **R** are the covariance matrices of the background and observation errors  
Choice of matrix is crucial to success of assimilation scheme

# Choosing the lengthscale for B

Increment (Analysis-Background) for assimilation of 1 waterline



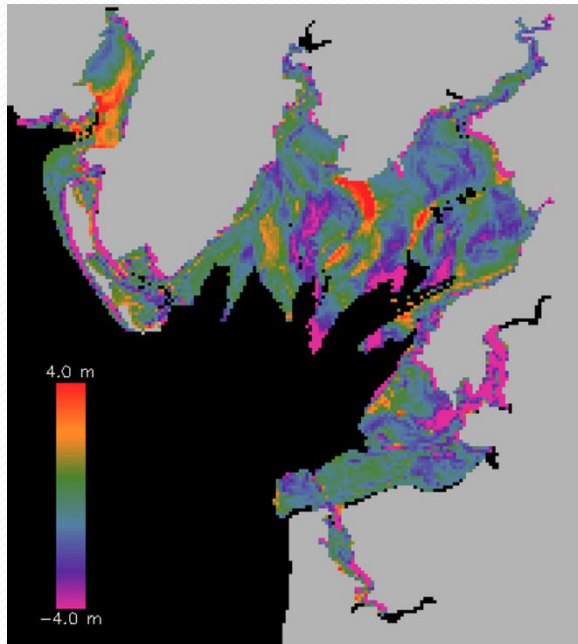
$L = 0.5$



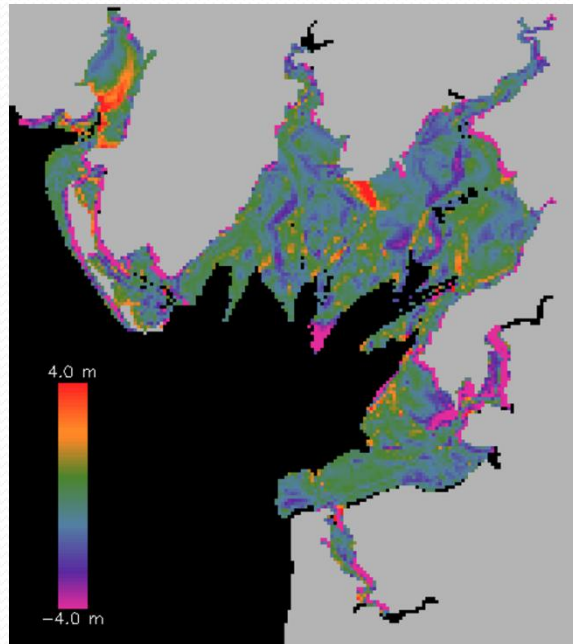
$L = 3.0$



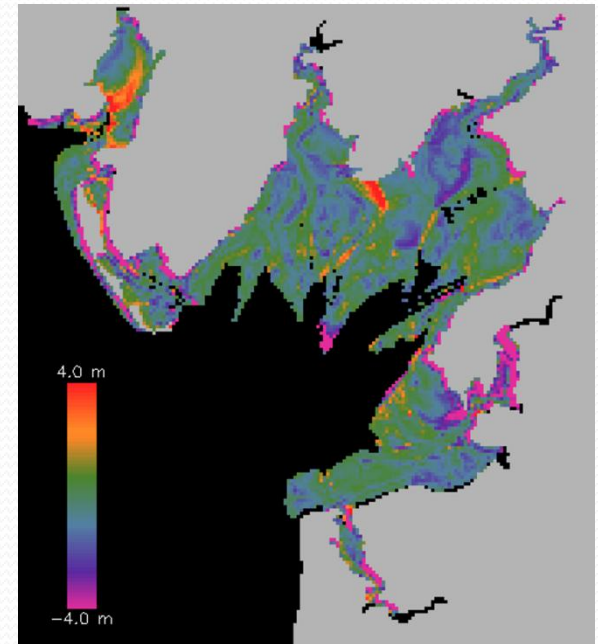
# Comparison with lidar validation data



No assimilation



With assimilation  
Calibrated params



With assimilation  
Weighted ensemble mean  
Parameter ensemble

# Parameter estimation

- Numerical models suffer from errors in their initial conditions and parameters.
- Initial conditions are often estimated by data assimilation; combining model predictions with observational data to produce an updated model state (the analysis) whilst keeping the model parameters fixed.
- Even with perfect initial data, inaccurate model parameters will lead to the growth of prediction errors.
- Using *state augmentation* in data assimilation we can estimate model parameters concurrently with the state.

# Joint state-parameter estimation

- Model state & parameter evolution

$$\mathbf{z}_{k+1} = \mathbf{f}(\mathbf{z}_k, \mathbf{p})$$

$$\mathbf{p}_{k+1} = \mathbf{p}_k.$$

- Augmented system model

$$\mathbf{w}_{k+1} = \begin{pmatrix} \mathbf{z}_{k+1} \\ \mathbf{p}_{k+1} \end{pmatrix} = \begin{pmatrix} \mathbf{f}(\mathbf{z}_k, \mathbf{p}_k) \\ \mathbf{p}_k \end{pmatrix} = \tilde{\mathbf{f}}(\mathbf{w}_k)$$

- Observations

$$\mathbf{y}_k = \tilde{\mathbf{h}}(\mathbf{w}_k) + \boldsymbol{\delta}_k = \tilde{\mathbf{h}} \begin{pmatrix} \mathbf{z} \\ \mathbf{p} \end{pmatrix} + \boldsymbol{\delta}_k = \mathbf{h}(\mathbf{z}_k) + \boldsymbol{\delta}_k$$

# Background error covariance

$$J(\mathbf{w}_k) = (\mathbf{w}_k - \mathbf{w}_k^b)^T \mathbf{B}_k^{-1} (\mathbf{w}_k - \mathbf{w}_k^b) + (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))^T \mathbf{R}_k^{-1} (\mathbf{y}_k - \tilde{\mathbf{h}}_k(\mathbf{w}_k))$$

background term      observation term

State background error covariance      state-parameter cross covariance

$$\mathbf{B}_k = \begin{pmatrix} \mathbf{B}_{zz_k} & \mathbf{B}_{zp_k} \\ (\mathbf{B}_{zp_k})^T & \mathbf{B}_{pp_k} \end{pmatrix}$$

parameter background error covariance

For joint state-parameter estimation, it is important that the a priori cross-covariances between the parameters and the state are well specified.

# A hybrid approach

Combines ideas from 3D-Var and the extended Kalman filter (EKF)

- assumes  $\mathbf{B}_{zz}$  and  $\mathbf{B}_{pp}$  fixed
- uses a flow dependent state-parameter cross covariance  $\mathbf{B}_{zp_k}$

$$\mathbf{B}_{k+1} = \begin{pmatrix} \mathbf{B}_{zz} & \mathbf{N}_k \mathbf{B}_{pp} \\ \mathbf{B}_{pp} \mathbf{N}_k^T & \mathbf{B}_{pp} \end{pmatrix}$$

where  $\mathbf{N}_k = \left. \frac{\partial \mathbf{f}(\mathbf{z}, \mathbf{p})}{\partial \mathbf{p}} \right|_{\mathbf{z}_k^a, \mathbf{p}_k^a}$



# A simple sediment transport model

Based on the sediment conservation equation

$$\frac{\partial z}{\partial t} = - \left( \frac{1}{1 - \varepsilon} \right) \frac{\partial q}{\partial x} \quad \text{with} \quad q = Au^n$$

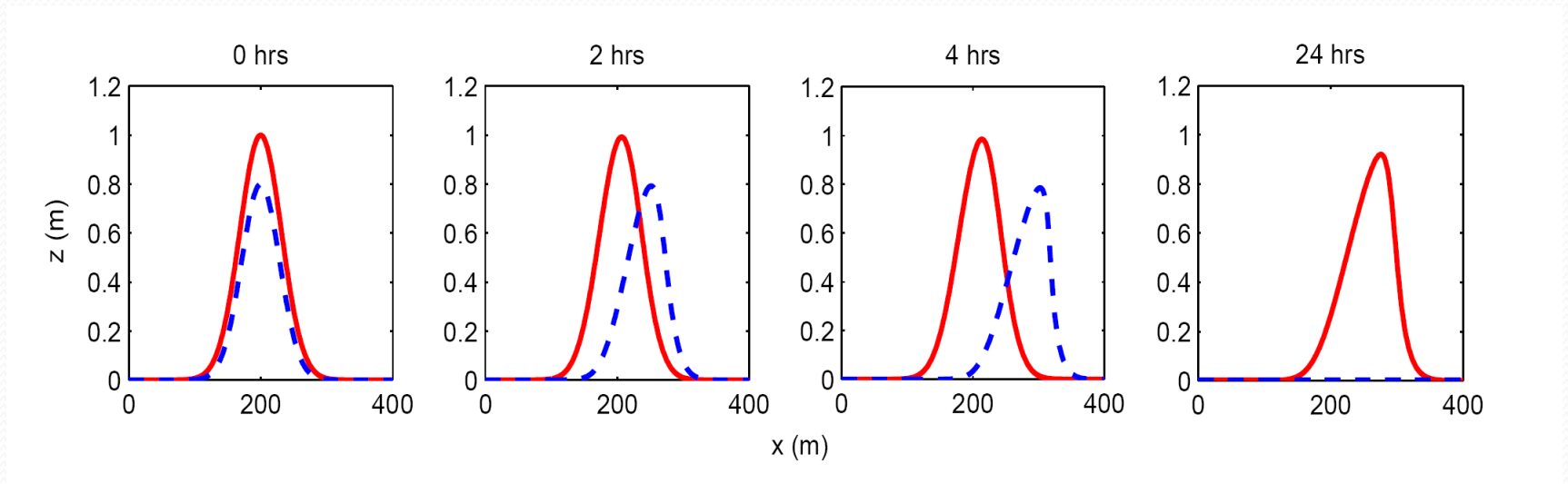
where  $z(x,t)$  is the bathymetry,  $t$  is time,  $\varepsilon$  is the sediment porosity,  $q$  is the sediment transport rate,  $u(x,t)$  is the depth averaged current and  $A$  and  $n$  are constant parameters.

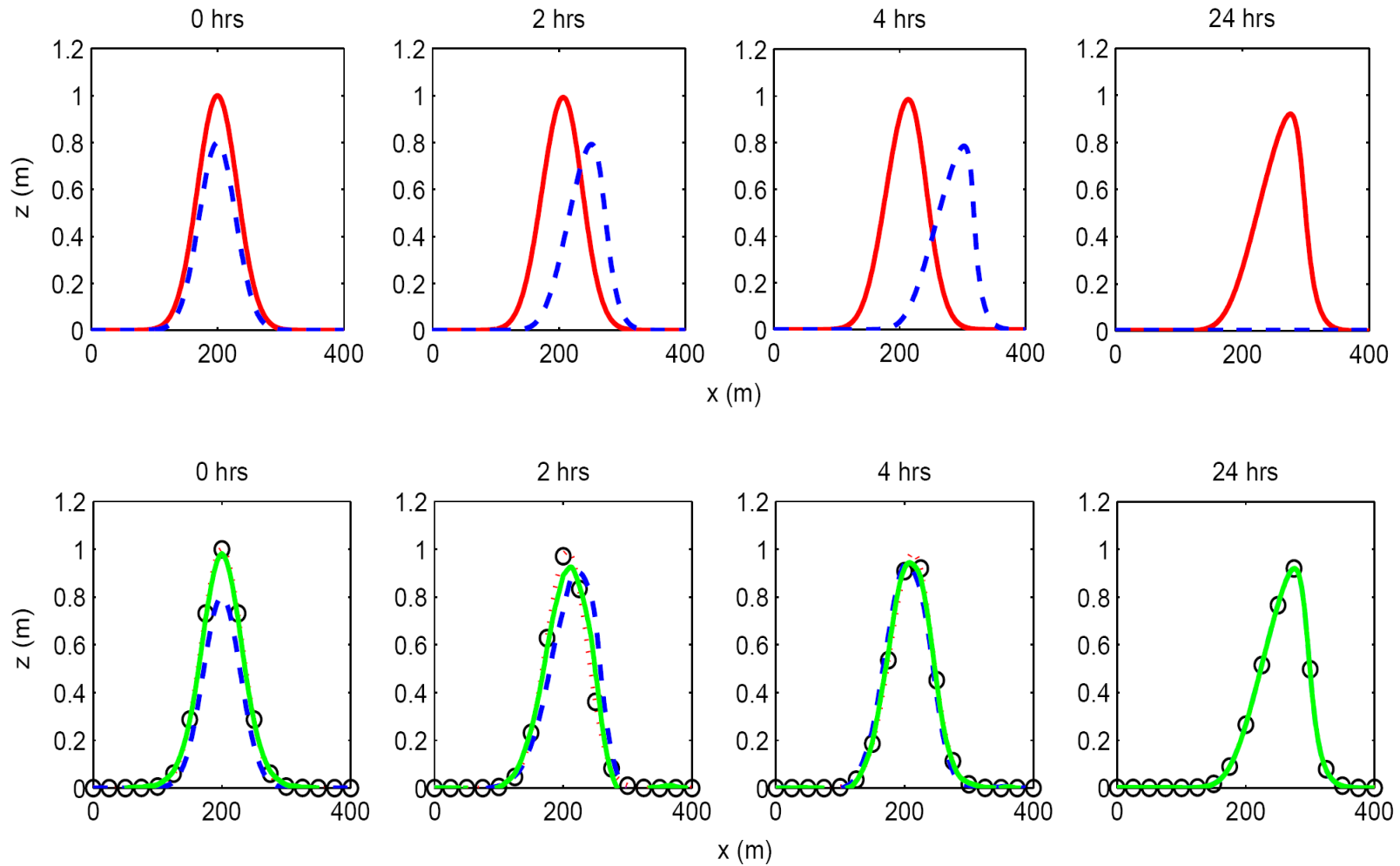
Can we use data assimilation to estimate the parameters  $A$  and  $n$ ?

# Experiments

- Identical twin:
  - reference solution generated using ‘true’ parameter values  $A = 0.002 \text{ ms}^{-1}$  and  $n = 3.4$
  - model then re-run with incorrect initial bathymetry and parameter values
- Observations assimilated sequentially at regular time intervals
  - taken from reference solution & assumed perfect
  - the 3D-Var cost function is minimized iteratively using a quasi-Newton descent algorithm
- Background error covariances
  - $\mathbf{B}_{zz}$  fixed
  - $\mathbf{B}_{zpk}$  time varying

# without data assimilation ...

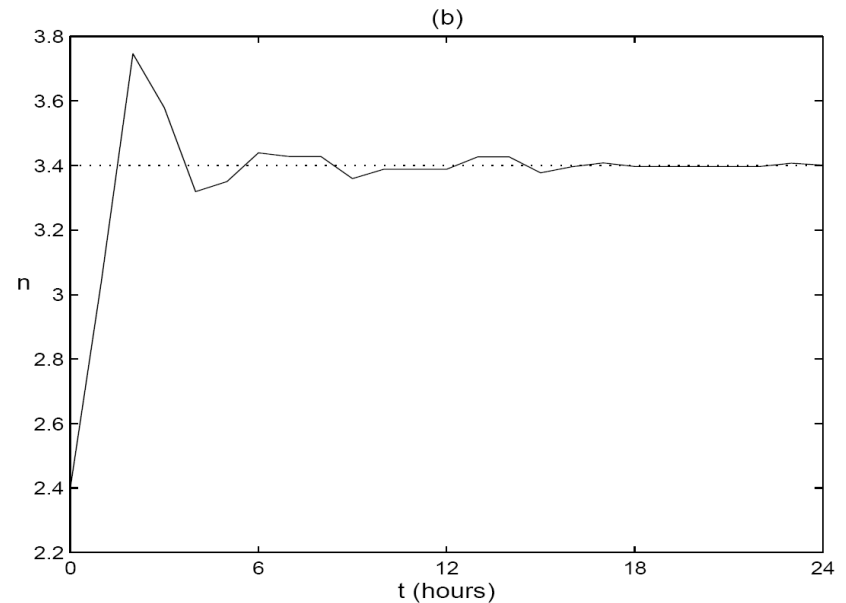
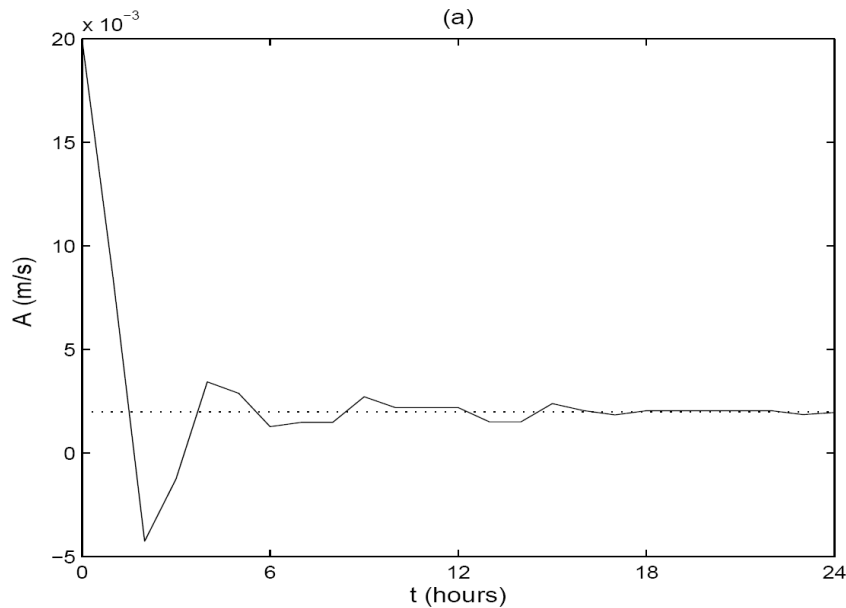




with data assimilation

# Parameter estimates

Initial estimates (a)  $A_o = 0.02 \text{ ms}^{-1}$  (b)  $n_o = 2.4$



$(A_{\text{true}} = 0.002 \text{ ms}^{-1}, n_{\text{true}} = 3.4)$



# Summary

- Up-to-date knowledge of near-shore coastal bathymetry is important in flood prediction and risk management
- Assimilated SAR waterline data into a model of Morecambe Bay to keep the model on track
- Best results are obtained using an ensemble of parameters
- Developed a new hybrid data assimilation scheme for joint estimation of model parameters and state.
- Recovers the true parameter values to a good level of accuracy, even when observations are noisy.
- Relatively simple to implement and computationally inexpensive to run
- Method also successfully applied to a range of simple dynamical system models.
- Expect this new technique to be easily transferable to more complex models.